

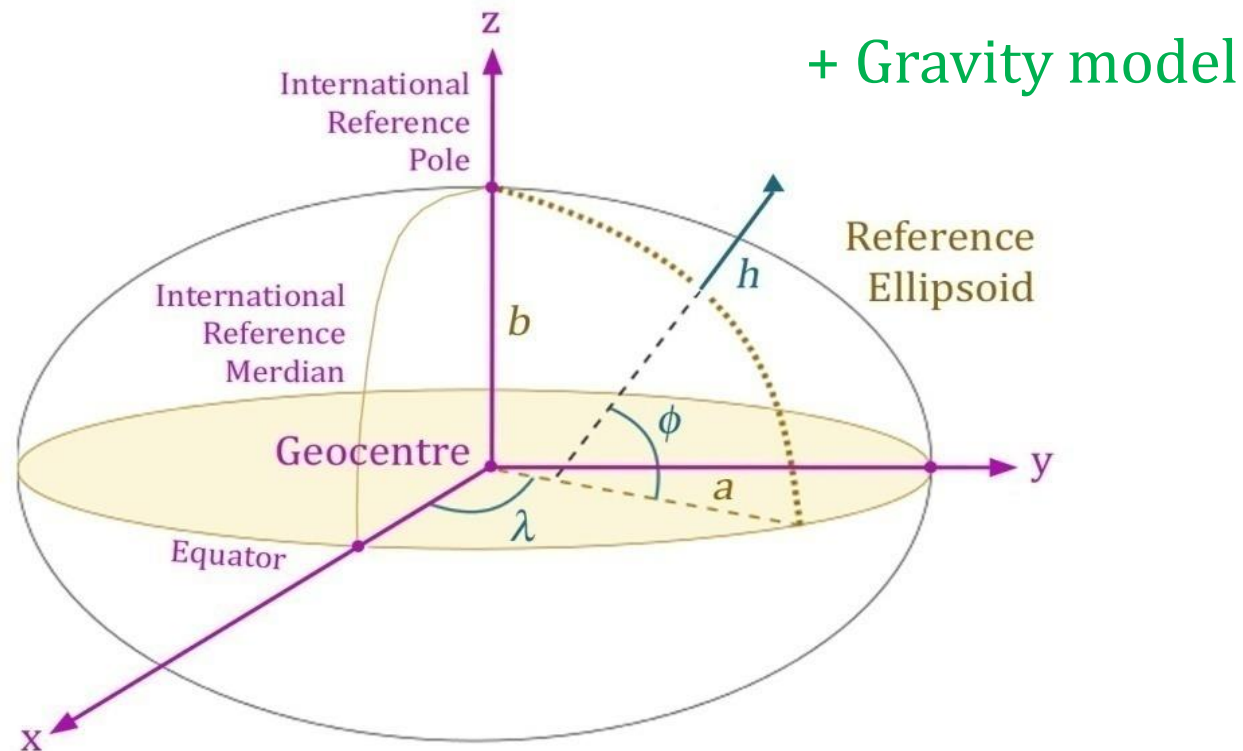
Reference Frames, Transformations and GIS

Richard Stanaway

School of Civil and Environmental Engineering

University of New South Wales

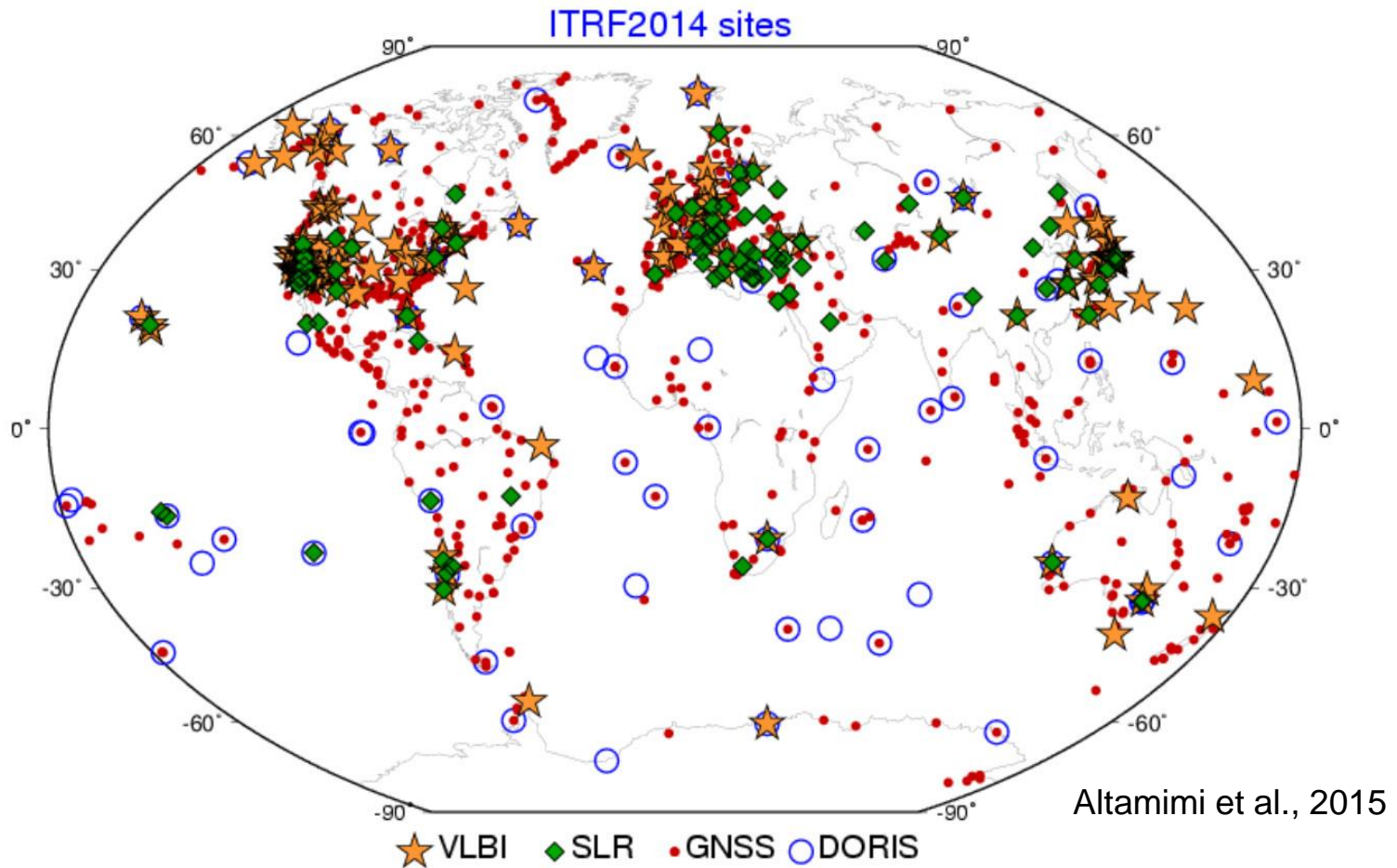
Terrestrial Reference System



In principle a TRS should be invariant with time

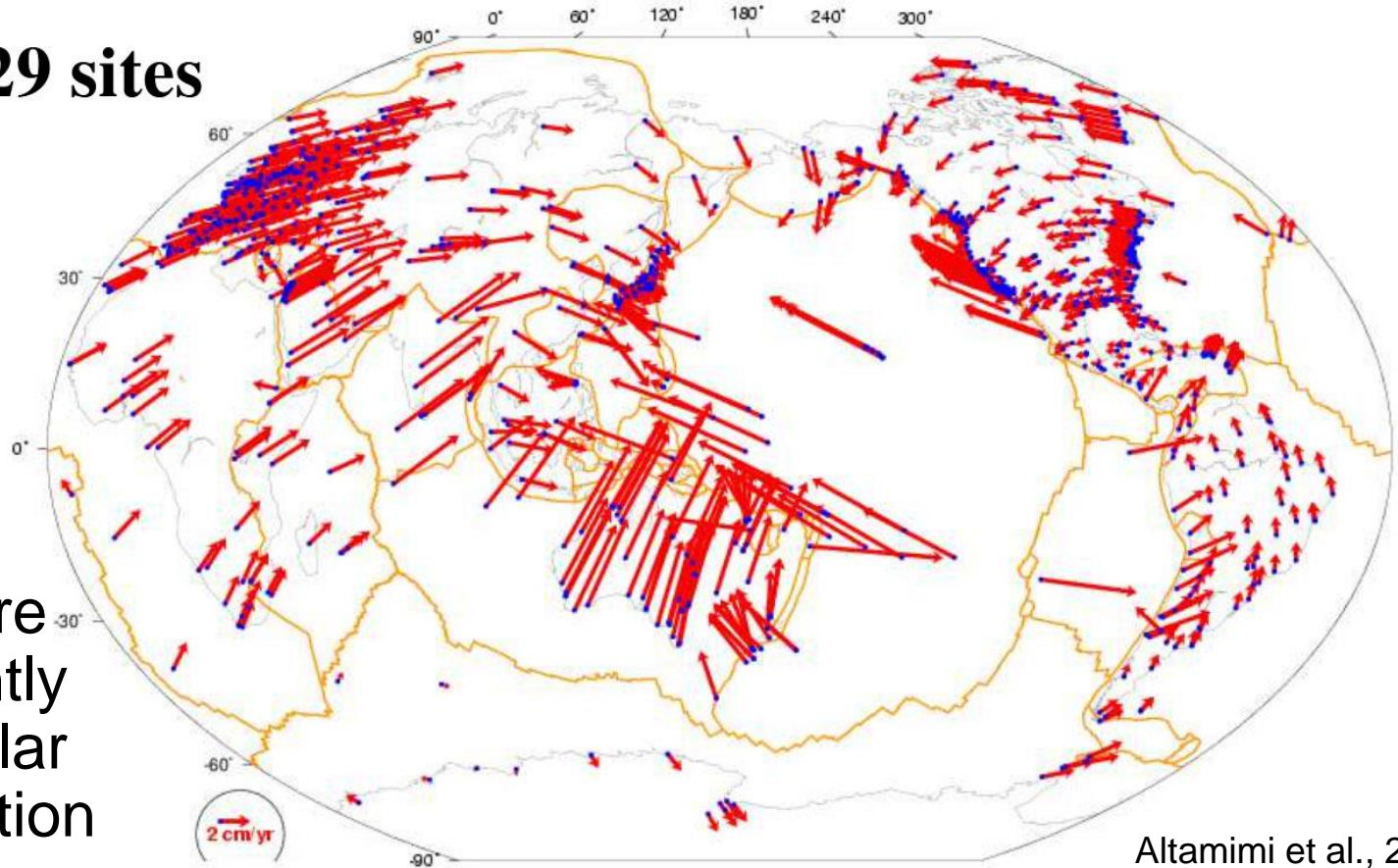
example: Geocentric ITRS

International Terrestrial Reference Frame



ITRF Kinematics – NNR site velocities

829 sites

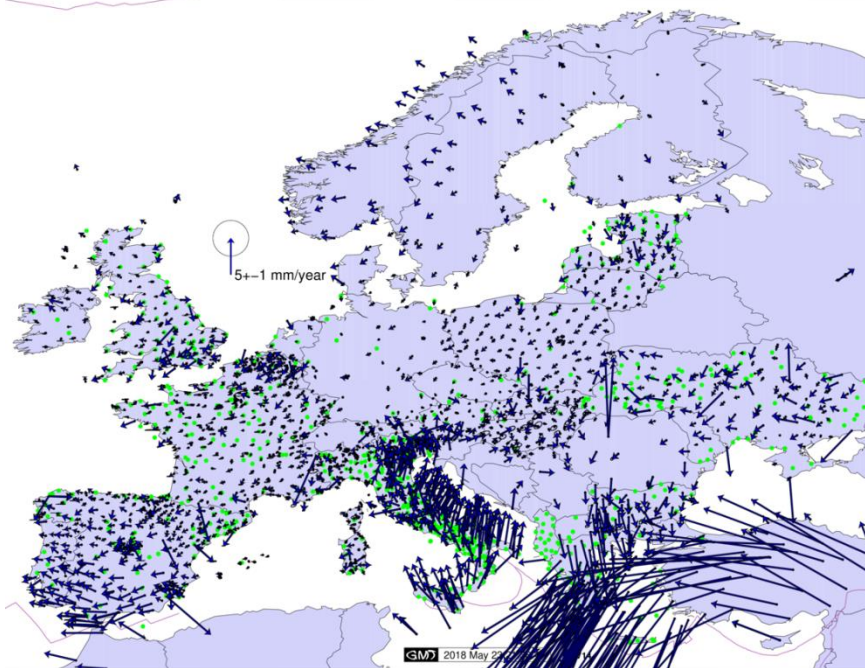


Velocities are predominantly due to secular tectonic motion and GIA

Altamimi et al., 2015

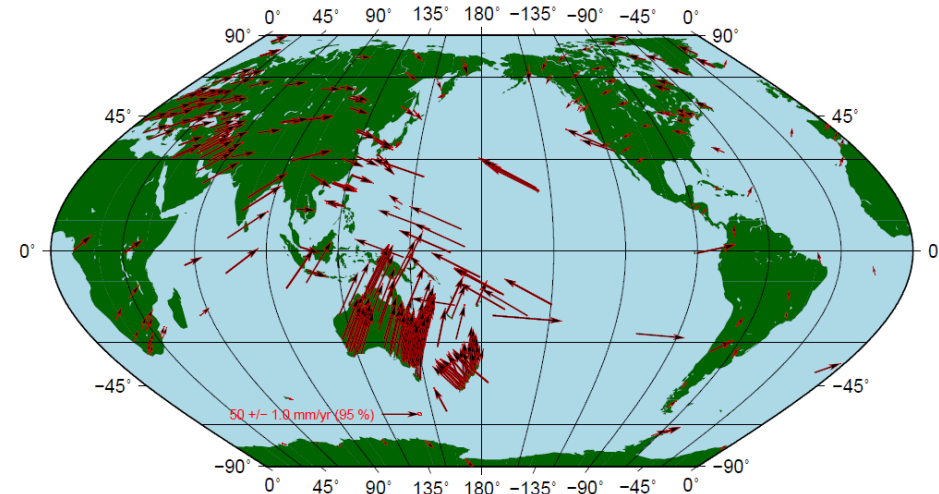
Regional Reference Frames

Plate fixed



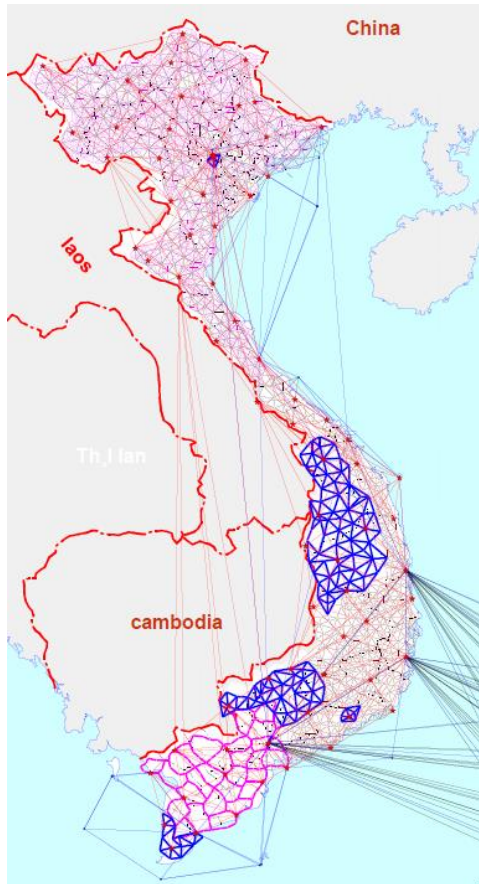
e.g. ETRF – velocities minimised
for most of Eurasian plate
(figure: EUREF, 2018)

NNR (No-Net-Rotation)



e.g. APREF – no dominant plate
within frame coverage so NNR
model is used (figure: Hu, 2014)

National & Local Reference Frames



Characteristics:

Now usually a fixed epoch of ITRF using GNSS
(recent geocentric frames)

e.g. VN2000 for mainland Vietnam
(figure: Vietnam Dept. of Surveying Mapping, 2016)

or astronomical determination of frame origin
(pre space geodetic era)
(frame often not geocentric)

e.g. Australian Geodetic
Datum 1966 (AGD66)
(figure: Paul Wise, 2016)



Static and time-dependent frames

Static (time-invariant) – no displacement is assumed

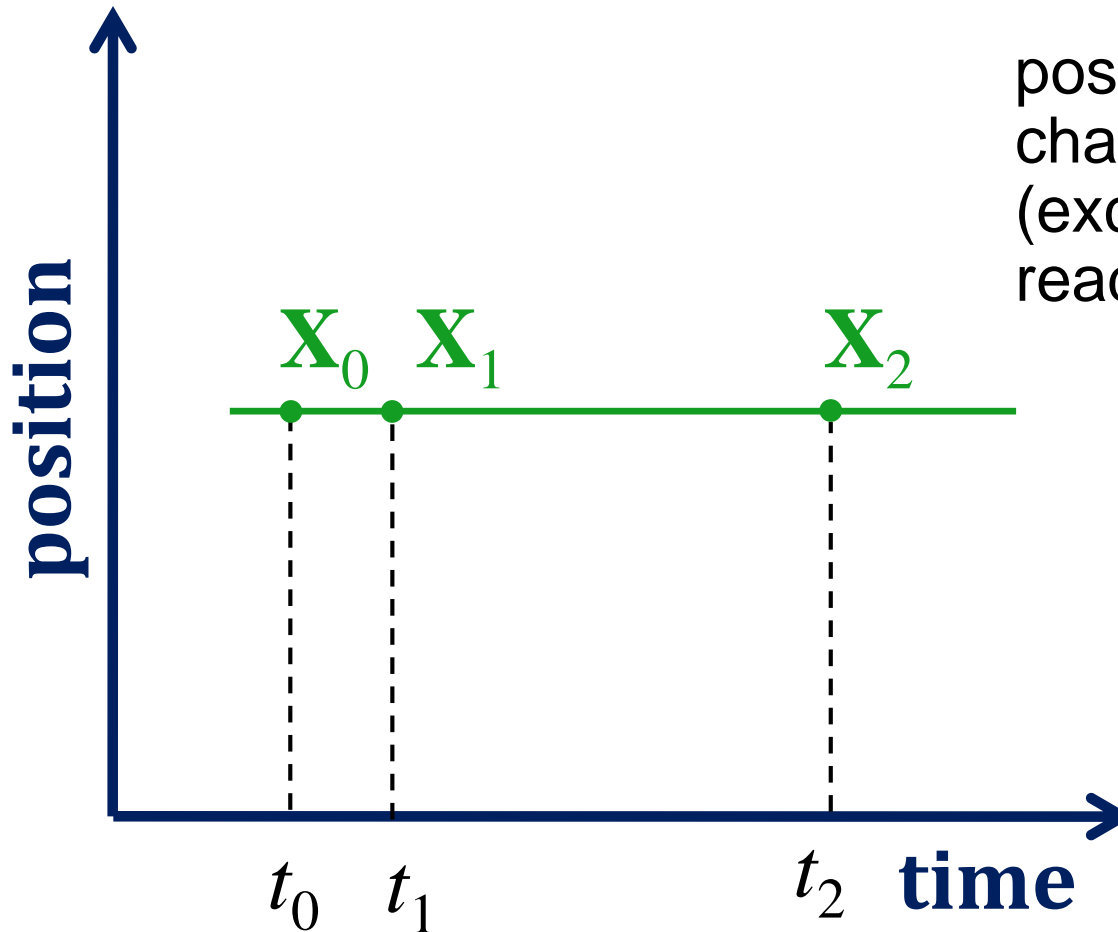
Kinematic (time dependent)

- includes all displacements wrt. ITRS
(e.g., tectonic motion, glacial isostatic adjustment,
coseismic deformation)

Semi-kinematic (time dependent)

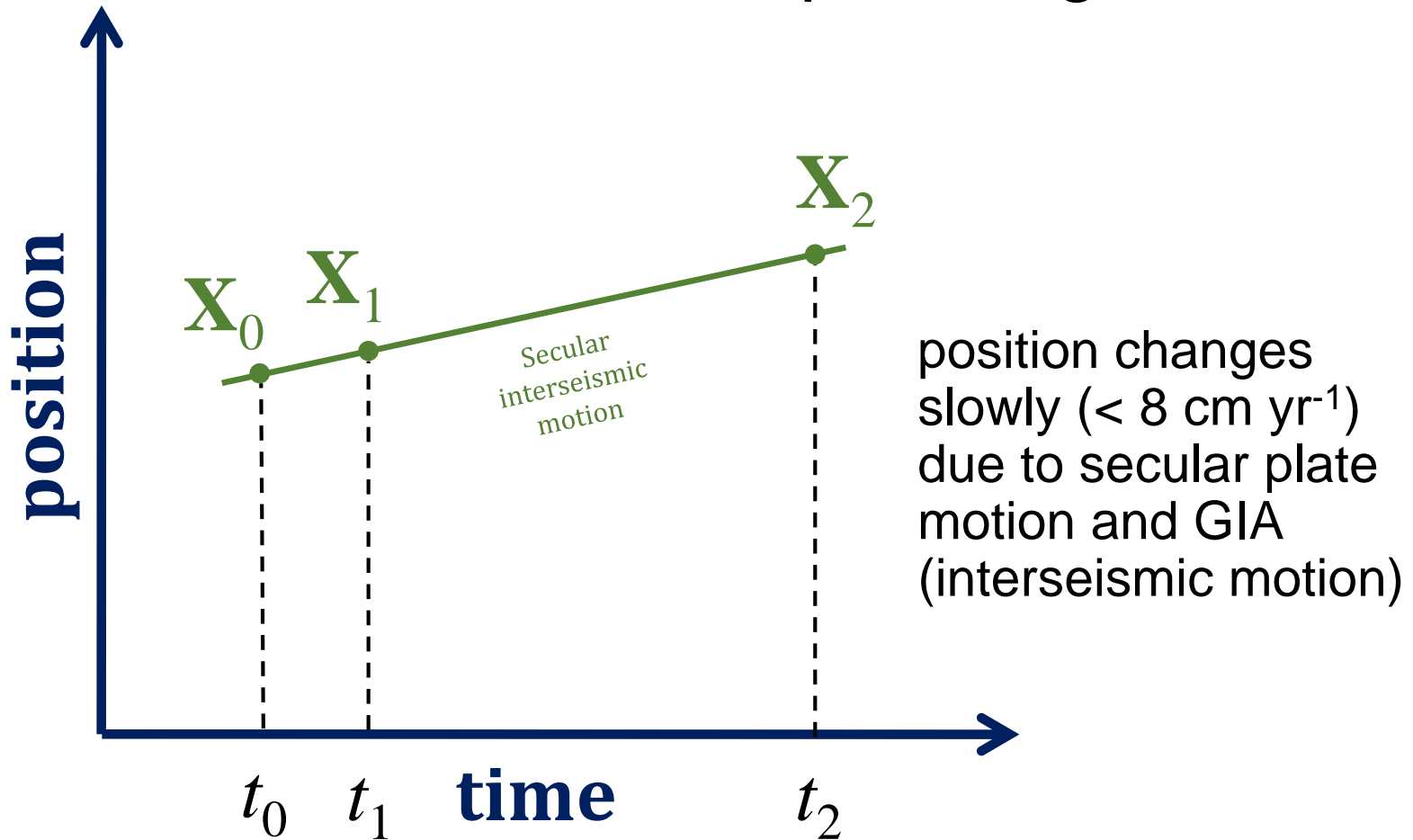
- secular tectonic motion is modelled out
- coseismic displacement models applied
after earthquakes

Static reference frame

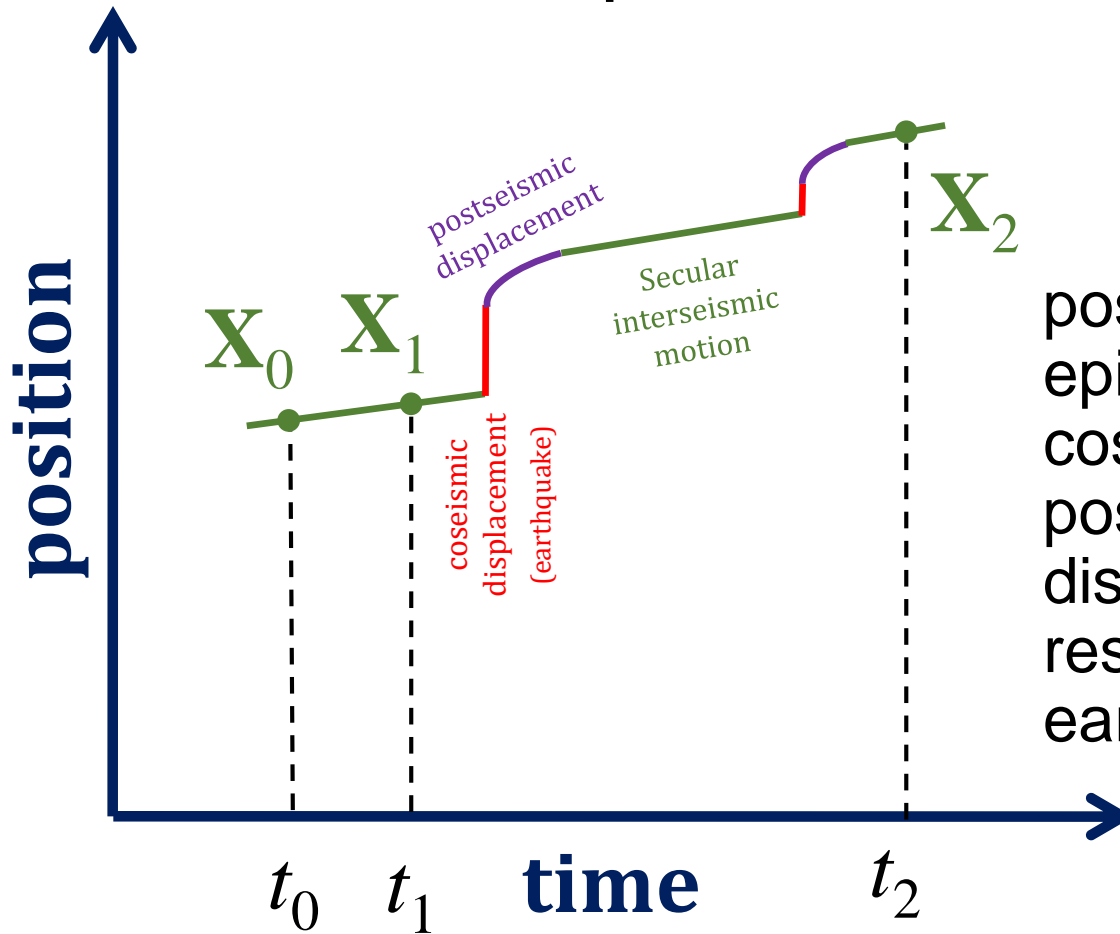


position does not
change with time
(except perhaps for
readjustments)

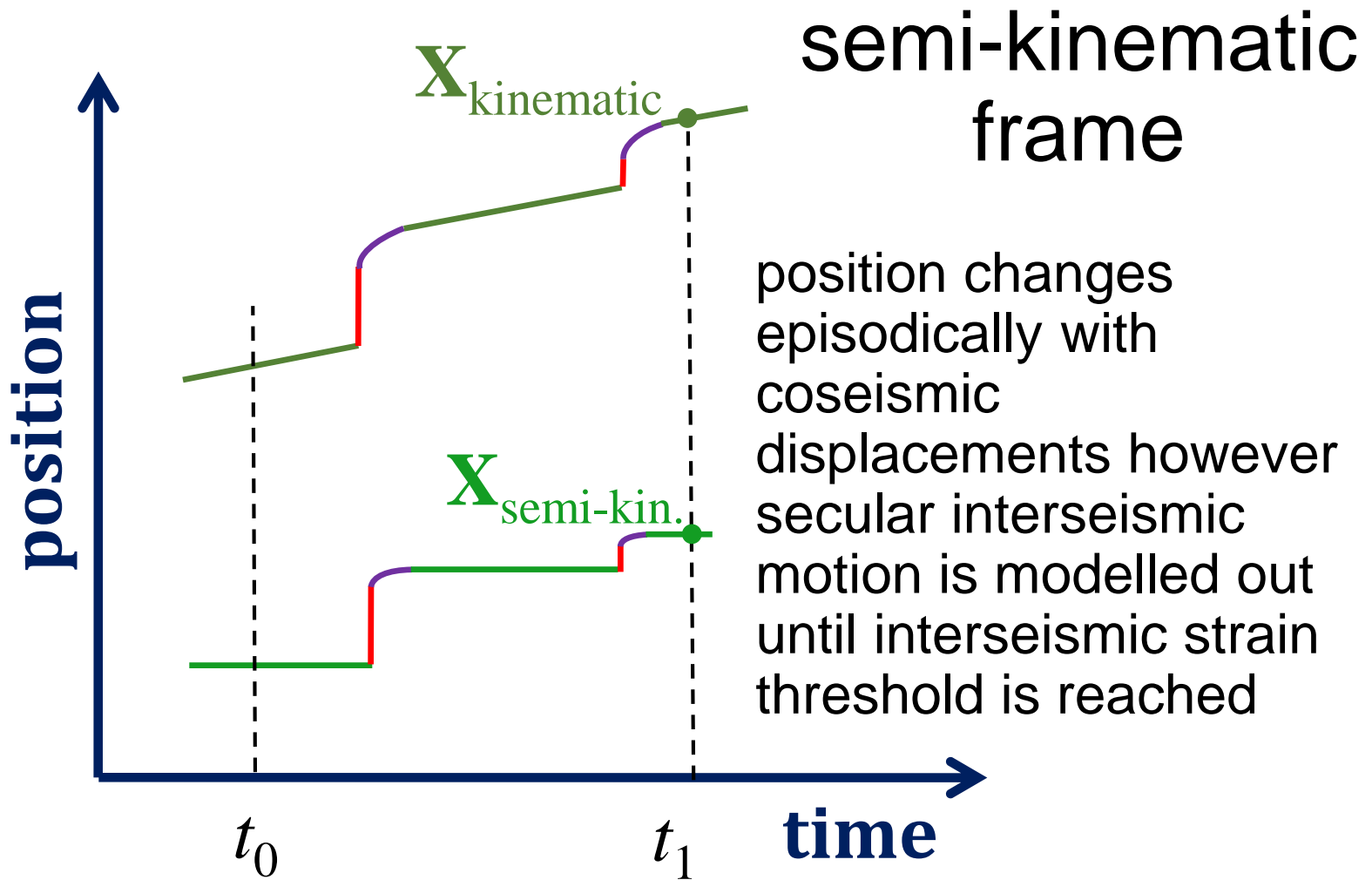
Kinematic frame – stable plate regions



Kinematic frame – plate boundary zones



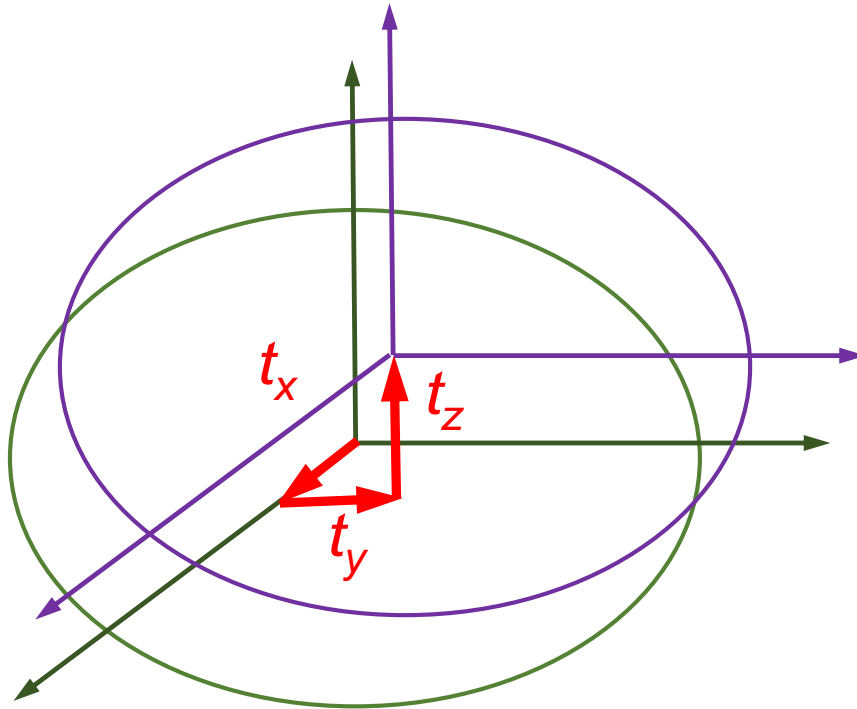
position changes episodically with coseismic and postseismic displacements resulting from earthquakes



Interframe transformations

	Static frame	time-dependent frame
Conformal Parametric	3 parameter 7 parameter	14 parameter (7 parameters at reference epoch + rates)
Grid model	displacement grid distortion grid	interseismic velocity model + coseismic displacement grid + postseismic decay term grid

3 parameter transformation



$$\mathbf{X}_B = \mathbf{X}_A + \mathbf{T}_{AB}$$

or

$$\begin{bmatrix} X \\ Y \\ Z \end{bmatrix}_B = \begin{bmatrix} X \\ Y \\ Z \end{bmatrix}_A + \begin{bmatrix} t_x \\ t_y \\ t_z \end{bmatrix}_{AB}$$

usually sufficient for most transformations between
homogeneous geocentric RF (no rotation or scale difference)

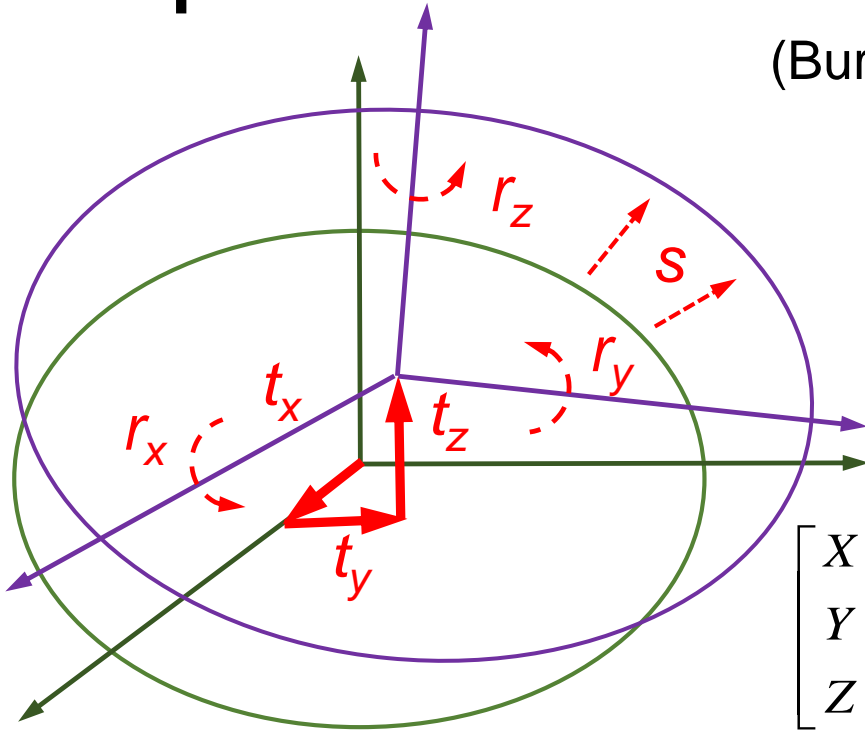
7 parameter transformation

(Bursa-Wolf simplification)

$$\mathbf{X}_B = \mathbf{T}_{AB} + (1 + s)\mathbf{R}_{AB}\mathbf{X}_A$$

or

$$\begin{bmatrix} X \\ Y \\ Z \end{bmatrix}_B = \begin{bmatrix} t_x \\ t_y \\ t_z \end{bmatrix}_{AB} + (1 + s) \begin{bmatrix} 1 & r_z & -r_y \\ -r_z & 1 & r_x \\ r_y & -r_x & 1 \end{bmatrix}_{AB} \begin{bmatrix} X \\ Y \\ Z \end{bmatrix}_A$$



used for most classical RF transformations (e.g., between a geocentric and astronomical RF). Bursa-Wolf method assumes small rotations (<10").

Coordinate Frame (CF) convention illustrated.

Position Vector (PV) rotations are defined in opposite sense

14 parameter transformation

7-parameter transformation at reference epoch t_0

+ parameter rates to estimate 7-parameters at different epochs

Used for static-kinematic or kinematic-kinematic RF transformations

model is reasonably good within stable tectonic plates

Plate Motion Model

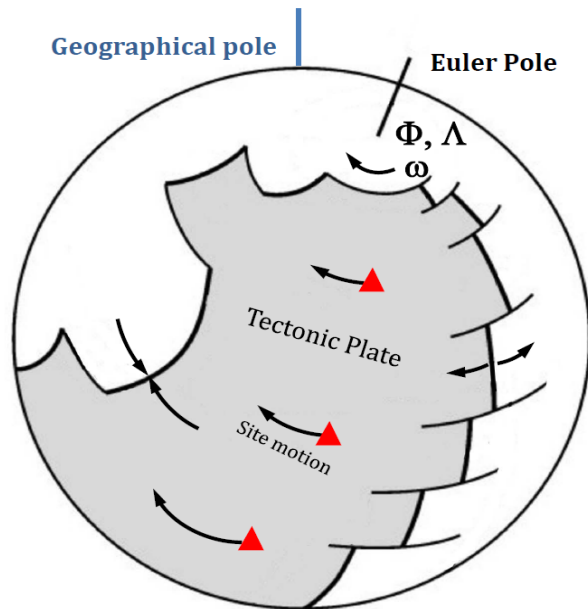


Plate	Euler pole of rotation			Equivalent Cartesian angular velocity		
	Φ ($^{\circ}$)	Λ ($^{\circ}$)	ω ($^{\circ}$ Ma $^{-1}$)	Ω_x (Rad Ma $^{-1}$)	Ω_y (Rad Ma $^{-1}$)	Ω_z (Rad Ma $^{-1}$)
Antarctic	58.8	-127.4	0.219	-0.001202	-0.001571	0.003272
Arabian	51.2	-6.7	0.515	0.005595	-0.000659	0.007001
Australian	32.4	38.1	0.631	0.007321	0.005730	0.005890
Eurasian	55.1	-99.1	0.261	-0.000412	-0.002574	0.003733
Indian	51.6	-0.2	0.516	0.005595	-0.000024	0.007049
Nazca	45.8	-102.2	0.629	-0.001614	-0.007486	0.007869
North American	-5.2	-88.0	0.194	0.000116	-0.003365	-0.000305
Nubian	49.7	-80.8	0.267	0.000480	-0.002977	0.003554
Pacific	-62.6	111.3	0.679	-0.001983	0.005076	-0.010516
South American	-19.1	-131.9	0.119	-0.001309	-0.001459	-0.000679
Somalian	47.7	-98.7	0.332	-0.000587	-0.003849	0.004286

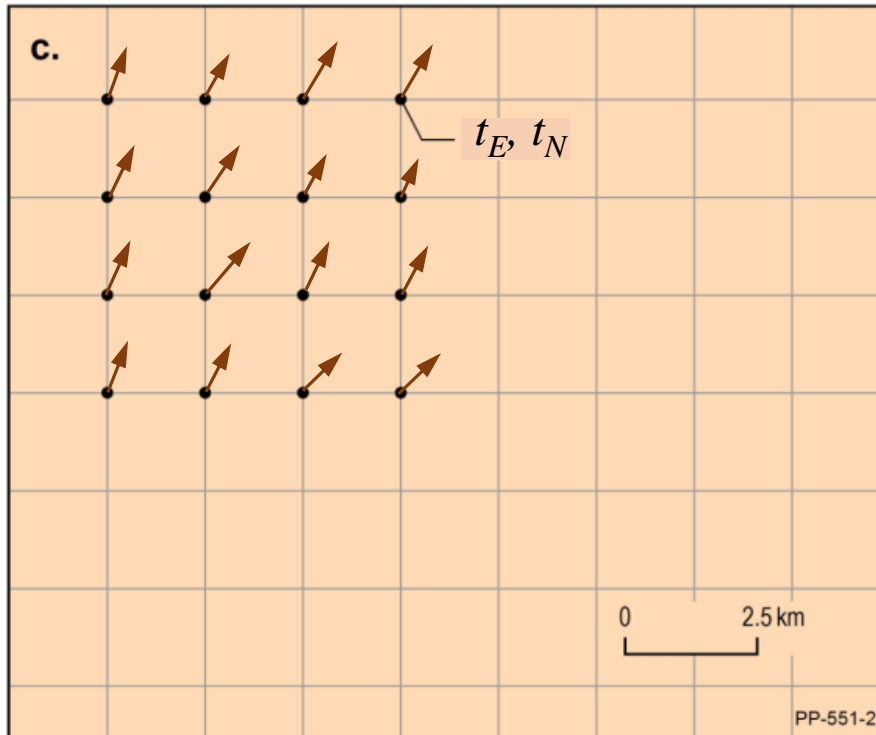
ITRF2014 Plate Motion Model (from Altamimi *et al.*, 2017)

Euler pole can be estimated by LS inversion of interseismic site velocities within a stable plate

$$\begin{aligned}\Omega_x &= \cos(\Phi) \cos(\Lambda) \omega \\ \Omega_y &= \cos(\Phi) \sin(\Lambda) \omega \\ \Omega_z &= \sin(\Phi) \omega\end{aligned}$$

Rotation rates can be used in a 14 parameter model with zeroes for other parameters (now used for GDA2020 to ITRF2014 transformations)

Displacement and Distortion Grids



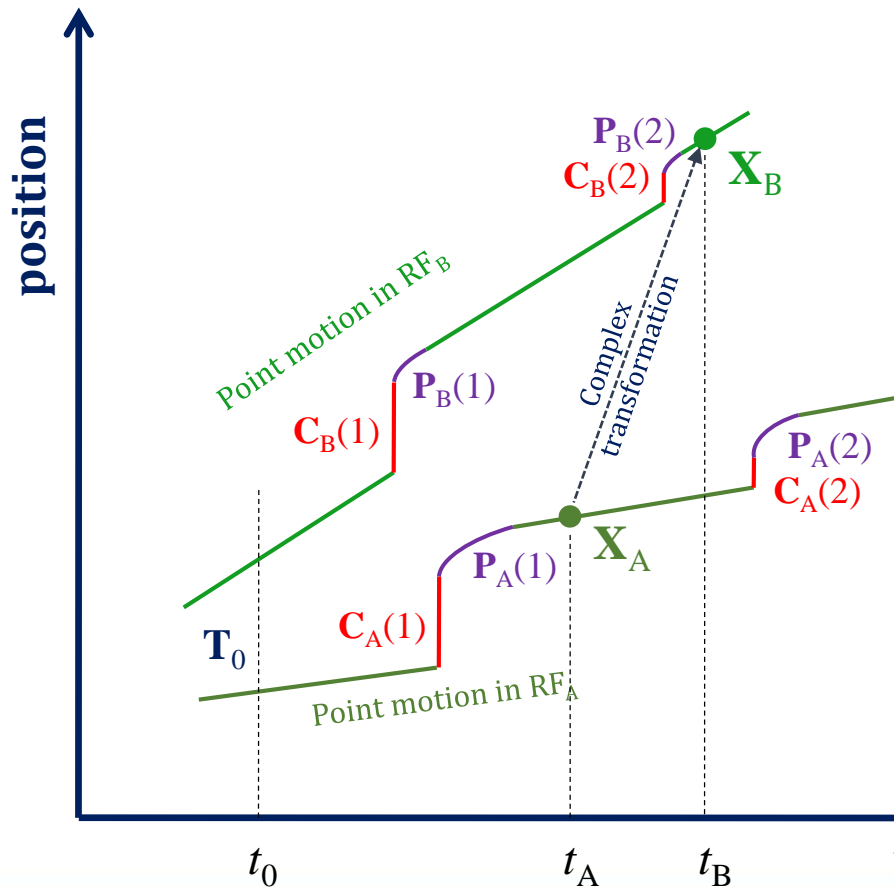
adapted from GDA2020 Technical Manual, ICSM, 2018

$$\begin{bmatrix} E \\ N \end{bmatrix}_B = \begin{bmatrix} E \\ N \end{bmatrix}_A + \begin{bmatrix} t_E \\ t_N \end{bmatrix}$$

Topocentric shifts estimated by
 bilinear interpolation of grid model
 (e.g., in NTv2 format)

Ideal for heterogeneous RF
 transformations and handling local
 distortions

Complex time-dependent transformations



$$\mathbf{X}_B = \mathbf{X}_A - \Delta\mathbf{X}_A + \mathbf{T}_0 + \Delta\mathbf{X}_B$$

where

$$\Delta\mathbf{X}_A = \mathbf{D}_A + \sum_{t_0 \geq C > t_A} \mathbf{C}_A + \sum_{t_0 \geq P > t_A} \mathbf{P}_A$$

and

$$\Delta\mathbf{X}_B = \mathbf{D}_B + \sum_{t_0 \geq C > t_B} \mathbf{C}_B + \sum_{t_0 \geq P > t_B} \mathbf{P}_B$$

- \mathbf{D}_A is secular displacement in RF_A
- \mathbf{D}_B is secular displacement in RF_B
- \mathbf{C}_A is coseismic displacement in RF_A
- \mathbf{C}_B is coseismic displacement in RF_B
- \mathbf{P}_A is postseismic displacement in RF_A
- \mathbf{P}_B is postseismic displacement in RF_B
- \mathbf{T}_0 is interframe translation

Topocentric frame

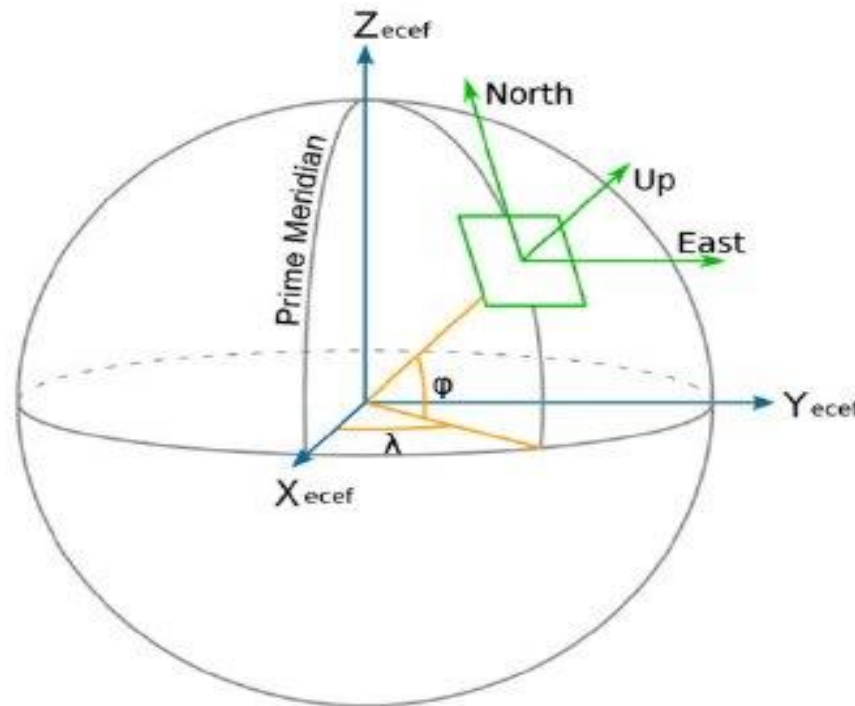


Image: Haasdyk and Janssen, 2011

Converting geocentric translation to topocentric (rate or shift)

$$\begin{bmatrix} \dot{E} \\ \dot{N} \\ \dot{U} \end{bmatrix} = \begin{bmatrix} -\sin \lambda & \cos \lambda & 0 \\ -\sin \phi \cos \lambda & -\sin \phi \sin \lambda & \cos \phi \\ \cos \phi \cos \lambda & \cos \phi \sin \lambda & \sin \phi \end{bmatrix} \begin{bmatrix} \dot{X} \\ \dot{Y} \\ \dot{Z} \end{bmatrix}$$

Converting topocentric translation to geocentric (rate or shift)

$$\begin{bmatrix} \dot{X} \\ \dot{Y} \\ \dot{Z} \end{bmatrix} = \begin{bmatrix} -\sin \lambda & -\sin \phi \cos \lambda & \cos \phi \cos \lambda \\ \cos \lambda & -\sin \phi \sin \lambda & \cos \phi \sin \lambda \\ 0 & \cos \phi & \sin \phi \end{bmatrix} \begin{bmatrix} \dot{E} \\ \dot{N} \\ \dot{U} \end{bmatrix}$$

A topocentric projection canvas is useful for complex transformation computations involving displacement grids

GIS transformations

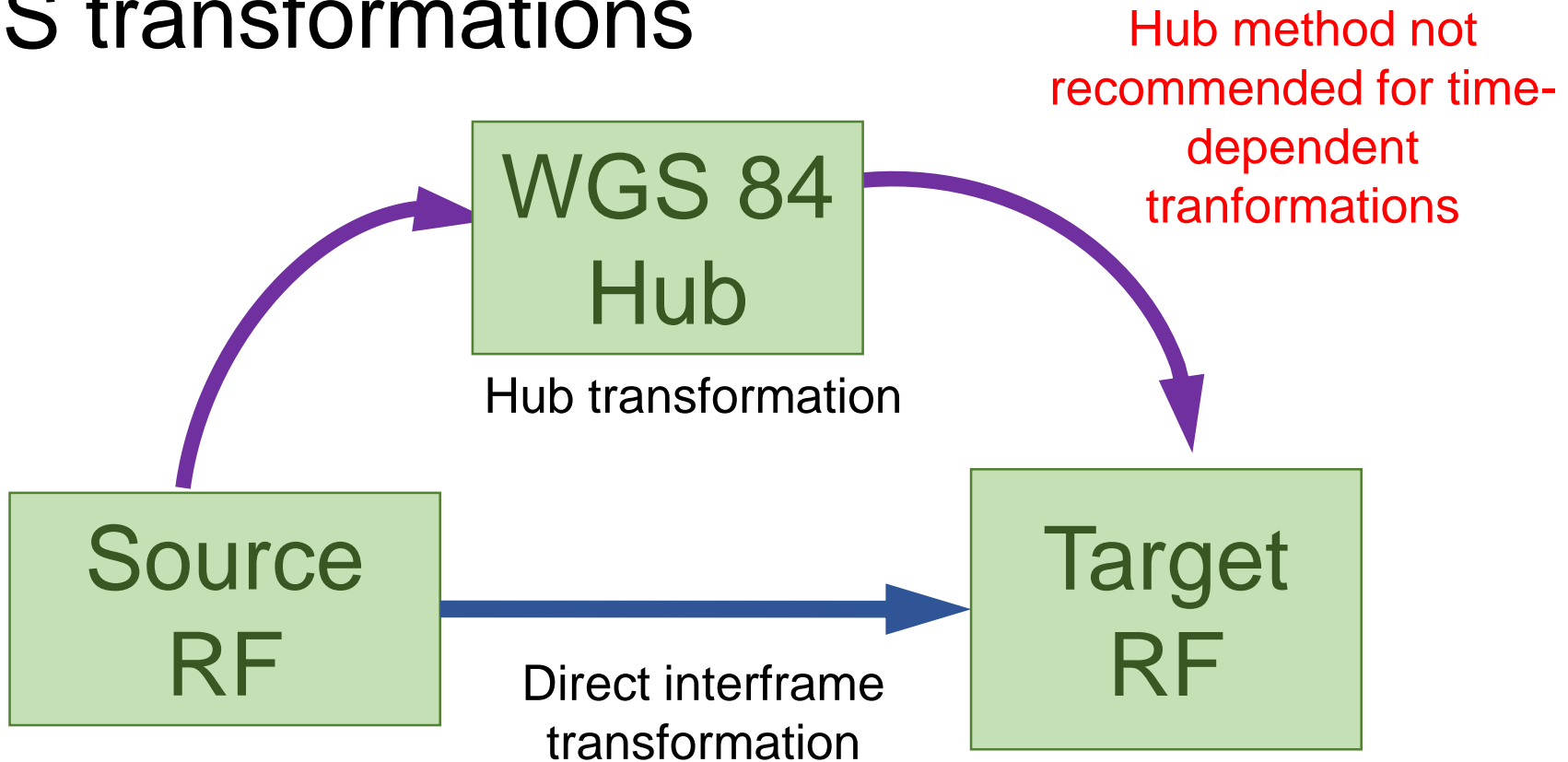
Main consideration is alignment of precise data defined in different RF and projections.

Geodetic Registries (EPSG and newly minted ISO TC211 Registry) standardise definitions and transformation workflows for use in GIS, however there are still limitations with complex time-dependent RF.

e.g. No standard for coseismic and postseismic displacement grids and no standardised epoch of WGS 84 when used as a hub transformation.

Direct interframe transformations are the preferred option.

GIS transformations



GIS Metadata Requirements

$\phi = 43^{\circ} 31' 32.3400''$ S
 $\lambda = 172^{\circ} 38' 23.4492''$ E
Geodetic NGZD2000 Epoch 2016.2

ϕ Precision: $\pm 0.0001''$
 λ Precision: $\pm 0.0002''$
Epoch: 2016.2
Type of position: Ellipsoidal
Reference Frame: ITRF96
Frame name: NGZD2000

X = -4593768.2707 m
Y = 593377.9433 m
Z = -4370031.2416 m
Cartesian NGZD2000 Epoch 2016.2

X Precision: ± 0.001 m
Y Precision: ± 0.002 m
Z Precision: ± 0.003 m
Epoch: 2016.2
Type of position: Cartesian
Reference Frame: ITRF96
Frame name: NGZD2000

ϕ, λ, h

X, Y, Z

GIS Metadata Requirements

$h = 23.126 \pm 0.007$ m Ellipsoidal NAD83(CSRS) Epoch 2013.2

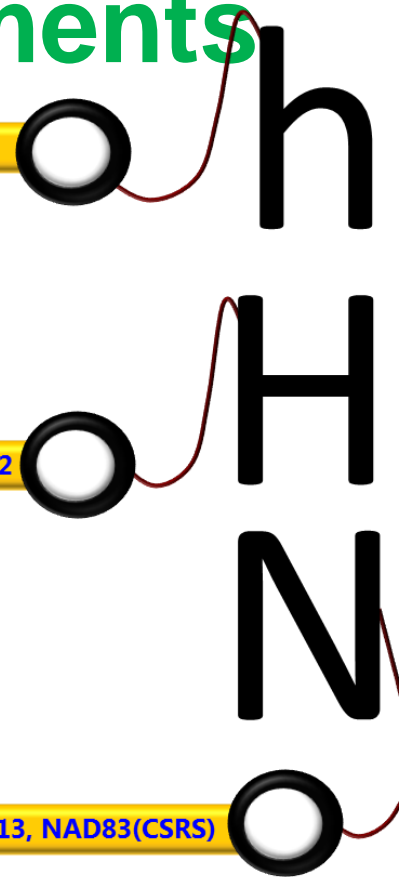
Height: 91.256 m
Precision: ± 0.007 m
Epoch: 2013.2
Type of height: Ellipsoidal
Height system: NAD83
Height frame name: CSRS

$H = 23.126 \pm 0.007$ m Orthometric CGVD2013(CG2013) Epoch 2013.2

Height: 101.61 m
Precision: ± 0.01 m
Epoch: 2013.2
Type of height: Orthometric
Height system: CGVD2013
Height frame: CGG2013

$N = -10.354 \pm 0.015$ m Geoidal CGG2013, NAD83(CSRS)

Geoid Height: -10.354 m
Precision: ± 0.015 m
Epoch: Static
Type of height: Geoidal
Model: CGG2013
Frame: NAD83(CSRS)



Thank you and Happy Easter in Viet Nam

Chúc mừng lễ Phục sinh

